Crustal Flow Pattern beneath the Tibetan Plateau
Constrained by Regional Lg-Wave Q Tomography

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Running title:
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Abstract

As a prominent geophysical anomaly, unusually high seismic wave attenuation is observed in the crust and upper mantle of the Tibetan Plateau, particularly along its northern area. Theoretical and laboratory investigations show that the strong seismic attenuation can indicate high temperatures and partial melting, which may decrease the viscosity of the material and cause it to flow. Thus, seismic attenuation distribution may provide useful constraints to the crust flows if they exist. Using Lg-wave Q tomography, we construct a broadband high-resolution crust attenuation model for the Tibetan Plateau and its surrounding regions. The maximum spatial resolution is approximately 1.0°×1.0° in well-covered areas and for frequencies between 0.05 and 1.5 Hz. Our broadband $Q_{\text{Lg}}$ model reveals there is an apparent low-$Q_{\text{Lg}}$ belt stretching along the northern and eastern Tibetan plateau. Combining the Lg-wave Q model with other geophysical data, two possible crust flow channels are found in the Tibetan Plateau. The main flow channel is from north to east and then turns to southeastern Tibet along the western edge of the rigid Sichuan basin, while a second channel starts from southern Tibet and crosses the Eastern Himalayan syntaxis.

Keywords

Lg attenuation; Q tomography; Lower-Crustal flow; Tibetan Plateau; Crustal deformation.
1. Introduction

A lower-crust flow model can explain many geological and geophysical observations in the Tibetan Plateau. Moreover, these observations provide important constraints to the dynamic processes in this region (Klemperer, 2006; Royden et al., 2008; Searle et al., 2011). Typically, the surface strain rates from the Global Positioning System (GPS) and earthquake data are consistent with a gravitationally driven flow model of a viscous lithosphere bounded by strong converging blocks in northern and southern Tibet (e.g., Clark and Royden, 2000; Flesch et al., 2001; Heidbach et al., 2010; Zhang et al., 2004). The low Pn velocities, inefficient Sn propagation, high Poisson’s ratios of approximately 0.35, and high seismic Lg-wave attenuation are observed in northern Tibet, suggesting that partial melting is existed within this region's crust (Fan and Lay, 2003b; Nelson et al., 1996; Owens and Zandt, 1997; Rodgers and Schwartz, 1998). The seismic and magnetotelluric observations revealed that, in eastern and southeastern Tibetan plateau, there are low-velocity and high-conductivity layers in the middle- and lower-crusts, which support a lower-crust flow model (e.g., Bai et al., 2010; Liu et al., 2006; Unsworth et al., 2005; Xu et al., 2007; Yao et al., 2008). However, whether the crust flow is widely spread throughout the entire Tibetan Plateau or is limited to certain narrow geological channels is still under debate. To explore the lower-crust flow pattern throughout the Tibetan Plateau, high-resolution regional measurements of the crust physical properties, such as the velocity, attenuation, anisotropy and electrical structures, are required (e.g., Fan and Lay, 2002; Li et al., 2008; Shapiro et al., 2004; Wang et al., 2013).

Seismic attenuation is usually an indicator of high temperatures and partial melts. An unusually high attenuation in both the crust and upper mantle is one of the first geophysical
anomalies discovered in Tibet, particularly in its northern region. The strong Pnl- and
Lg-wave attenuations are consistent with the strong Sn-wave attenuation in this region (Fan
and Lay, 2003b; Ni and Barazangi, 1983; Rodgers and Schwartz, 1998). Both Rodgers and
Schwartz (1998) and Fan and Lay (2003b) suggested that the strong attenuation results from
widespread partial melting in the northern Tibetan crust. Xie et al. (2004) found strong crustal
Lg-wave attenuation in the Yangbajing graben in southern Tibet and attributed the attenuation
to hydrothermal and magmatic fluid activities in the upper-crust. Based on deep seismic
sounding data from eastern Tibet, Wang et al. (2007) compared the amplitude difference of
seismic PmP waveforms between the observed and synthetic data. They suggested that the
weak PmP amplitudes resulted from a high attenuation in the lower crust and hence suggested
that a lower-crust flow is likely existed in this region. In the Tibetan Plateau, previous
attenuation studies are mostly limited within local regions or to very low resolutions because
of limited data (Bao et al., 2011; Fan and Lay, 2002, 2003a, b; McNamara et al., 1994;
Rodgers and Schwartz, 1998; Xie, 2002; Xie et al., 2004; Zhou et al., 2011). Until recently,
due to the lack of a high-resolution attenuation model for the Tibetan Plateau, it has been
difficult to link the attenuation information with the regional tectonics.

In this study, we develop a high-resolution Lg-wave attenuation model in the Tibetan
Plateau and its adjacent regions and investigate its connections to the thermal activities and
possible material movement in the lower crust and upper mantle.

2. Data and Methods

We collected 7545 broadband vertical-component digital seismograms recorded at 146
stations from 232 regional earthquakes between January 2001 and June 2008 with their ray paths penetrating the plateau. The waveforms were obtained from the China Earthquake Networks Center (CENC) and the Incorporated Research Institutions for Seismology (IRIS) consortium. The station parameters, including code, location, data resource, and affiliation, are listed in Tables S1 and S2 in the supplementary document. Both the CENC and the IRIS stations are equipped with broadband instruments having nearly flat velocity responses from 0.03 Hz to 8.0 Hz and one of the three sampling rates: 20, 40 and 50 points per second. The earthquake parameters are listed in Table S3 in the supplementary document. Shown in Fig. 1 is a topographic map overlapped with the main fault systems (light-blue lines), geotectonics (white lines), locations of the CENC (solid squares) and IRIS (triangles) stations, and epicenters of the earthquakes (crosses) used in this study. The waveforms were selected based on the criteria that these earthquakes were located in the crust, their magnitudes ranged between $m_b = 3.5$ and 6.0, and the epicenter distances were between 200 and 3000 km.

The data pre-processing was conducted following Zhao et al. (2010, 2013). We extracted the Lg-waveforms using a group-velocity window of 3.6-3.0 km/s and collected the noise time series in an equal-length window as the Lg phase before the first-arriving P wave. Then, we calculated Fourier spectra for both the Lg-wave and the noise, sampled the spectral amplitudes, and corrected for the noise effects. Our Lg-wave spectra calculation is illustrated in Fig. 2. In Fig. 2a, the solid and dashed lines denote the amplitude spectra of Lg and pre-P noise, where the circles and triangles denote the samples at 58 frequencies distributed log evenly between 0.05 and 10.0 Hz. From the signal and noise spectral amplitudes, we calculated the signal-to-noise ratios at individual frequencies (shown in Fig. 2b as solid
circles). A threshold of 2.0 is shown as a dashed line and was used for rejecting the low quality data. The noise-corrected Lg-wave spectrum is illustrated in Fig. 2c, where points below the threshold are dropped. After batch processing all regional waveforms, we obtained the source-station amplitudes at individual frequencies between 0.05 and 10.0 Hz. Following Xie et al. (2004) and Zhao et al. (2013), we extracted the dual-station data for individual frequencies from the source-station data. Both dual- and single-station data were used in the joint inversion for the Lg Q distribution and Lg-wave source functions (for details see Zhao et al., 2013). Using a checkerboard method (e.g., Zhao et al., 2013) with variable grid sizes from $0.8^\circ \times 0.8^\circ$ to $2^\circ \times 2^\circ$, we conducted resolution analyses for all 58 individual frequencies. Fig. 2d summarizes the quantities of available rays (for dual-station, single-station, and combined data sets) versus frequency, where the shaded areas illustrate the estimated resolutions for particular frequencies.

3. **Tomographic Model of Lg Attenuation**

Based on the above mentioned Lg dataset, we obtained a broadband attenuation model for the Tibetan Plateau and its surrounding regions, where $Q_{Lg}$ is distributed geographically as well as at 58 discrete frequencies between 0.05 and 10.0 Hz.

3.1 **$Q_{Lg}$ maps at Individual Frequencies**

Figs. 3a-c illustrate the $Q_{Lg}$ distributions at 0.5, 1.0, and 2.0 Hz, respectively, along with the major geological sutures (white lines) and active fault systems (thin black lines). Note that different color scales are used for these $Q_{Lg}$ images. The most prominent feature in these maps is that the high-frequency $Q_{Lg}$ is generally higher than the lower-frequency values. The lateral
$Q_{Lg}$ variations are consistent with regional tectonics. Compared to its surrounding areas, the Tibetan Plateau is characterized by strong $Lg$ wave attenuations, with low-$Q_{Lg}$ zones seen in the Songpan-Ganzi-Hoh xil (ST), Qiangtang (QT), and Lhasa Terrane (LT) regions, forming a belt first from west to east then turn to south. Shown in Fig. 3d are some 1 Hz $Q_{Lg}$ measurements from previous investigators (Fan and Lay, 2002, 2003a, b; Rai et al., 2009; Xie, 2002; Xie et al., 2004; Zhao et al., 2013). Comparison between Fig. 3b and 3d demonstrates that our result is consistent with the previous measurements but with higher resolution and better coverage. Shown in Figs. 3e and 3f are the ray-path coverage and the checkerboard resolution analyses at 1 Hz. In well-covered areas, the spatial resolution can reach to $1^\circ \times 1^\circ$ or higher.

There are strong $Q_{Lg}$ variations between different geology formations as well as within each unit. We investigated the $Lg$ attenuation in different geological formations by calculating their average values (Zhao et al., 2010, 2013). The geographically averaged $Q_{Lg}$ values versus frequency in selected tectonic regions are shown in Fig. 4. As an example, the light gray crosses in Fig. 4a are the inverted $Q_{Lg}$ values within the Songpan-Ganzi-Hoh xil terrane (ST). The result reveals the $Q_{Lg}$-frequency relationship but with large scatters. We calculated the mean $Q_{Lg}$ values within narrow frequency bands, and the results are shown as squares with error bars. These statistical results are more robust in characterizing the regional variations and frequency dependence of $Q_{Lg}$. Also labeled in Fig. 4a are the average $Q_0$ and its standard deviations. Fig. 4b summarizes the average $Q_{Lg}$ versus frequency relations for selected geological blocks, and these results are also listed in Table 1. The average $Q_{Lg}$ values between 0.2 and 2.0 Hz (shaded area in Fig. 4b) show larger regional variations, thus suitable for
characterizing the regional attenuation variations. The part of Tibetan Plateau with elevations above 4,000 m have an average $Q_0$ of 280 (194 – 406), much lower than the values of 374 (273-512) and 414 (232- 739) from North China Craton (NCC) and Northeast China (NEC) (Zhao et al., 2010, 2013). The regions surrounding the Tibetan plateau are mostly characterized by high $Q_0$ values (Tarim basin: 433, Altyn Tagh mountain: 517, Qaidam basin: 385, Alashan uplift: 452, Yinshan mountains: 444, Ordos: 395, and Sichuan basins: 456) except for Yungui Plateau, which has a relatively low $Q_0$ of 247.

3.2 Broadband $Q_{Lg}$ images

To explore the relationship between the broadband $Q_{Lg}$ and the regional geology, we use cross-sections to show the frequency dependence of the attenuation. Shown in Fig. 5 are six east-west cross-sections located at selected latitudes from north to south. The left column compares the $Q_{Lg}$ (0.2-2.0 Hz) (average $Q_{Lg}$ between 0.2 and 2.0 Hz), surface topography, and Moho depths from CRUST2.0 (Bassin et al., 2000). Illustrated in the right column are $Q_{Lg}$ versus frequency along these sections. For these cross sections, their latitudes are labeled in the left column, and the longitudes are indicated along the top and bottom. Also labeled in these sections are names of major geology blocks and the locations the lowest $Q_{Lg}$ appears (with arrows). Located in the northernmost portion of the Tibetan Plateau, the section along the 37.5°N latitude passes sequentially through the Tarim basin (TB), Eastern Kunlunshan terrane (EKT), Qaidam basin (QB), Qilianshan mountains (QM), and Ordos basin (OB). As shown in Fig. 5a, there are two apparent low-Q regions, EKT and QM, corresponding to mountain areas, while the three stable basins have relatively higher $Q_{Lg}$ values (as can be seen
from Fig. 3). The 35°N latitude section crosses the Qiangtang (QT), Songpan-Ganzi-Hoh xil (ST), and Eastern Kunlunshan terranes (EKT). Located in the North Tibetan Plateau, the crust in this region was suggested by Owens and Zandt (1997) to be partially melted due to high temperatures, thus having relatively low $Q_{Lg}$. In Figs. 5c and 5d, low-$Q_{Lg}$ anomalies occur near 89°E and 95°E, which agrees with the findings by Fan and Lay (2002, 2003a, b). The 32.5°N latitude section traverses the Qiangtang (QT) and Songpan-Ganzi-Hoh xil terranes (ST) and enters into the Sichuan basins (SB). Relatively low-$Q_{Lg}$ values are observed in eastern Tibet as shown by arrows in Fig. 5e. Located in the south Tibetan Plateau, section along the 30°N latitude sequentially passes through the Himalaya (HM), Lhasa (LT), Qiangtang (QT), and Songpan-Ganzi-Hoh xil terranes (ST) and the Sichuan basins (SB). In Figs. 5g and 5h, sections along 27.5°N latitude show apparently strong attenuation regions between 90° and 95°E in southern Tibet and near 100°E in southeast Tibet. Geophysical anomalies such as middle-crustal low-velocity, low-$Q_{Lg}$, high-heat flows and crustal electrical conductivities have been observed in southern Tibet (e.g., Langille et al., 2010; Wei et al., 2001; Xie et al., 2004; Zhang et al., 2011). The partial melting, resulting from collisional crustal thickening, is thought to be responsible for these geological anomalies (e.g., Beaumont et al., 2004; Nelson et al., 1996). The southeastern Tibet is also characterized by strong attenuation in the crust, as shown in Figs. 5g-l. It is commonly accepted that this region is an exit of the lower crust channel flow in eastern Tibet (e.g., Clark and Royden, 2000; Royden et al., 2008; Wang et al., 2013; Zhang et al., 2010).

4. The Possible Crustal Flow Pattern Constrained by $Lg$ Attenuation
The partially molten layer in the middle or lower crust is generated by the continental crust thickening and behaves like a fluid over the formation time scale of the Tibetan Plateau (e.g., Nelson et al., 1996). It is expected that the crustal flow is characterized by low viscosity, high temperature and partial melting. Several high-attenuation regions are observed in the crust and upper mantle in the Tibetan Plateau. Theoretical models and laboratory measurements show that strong seismic attenuation is usually an indicator of high temperatures and partial melts. Therefore, we try to use regional Lg-wave attenuation distributions to constrain the possible Tibetan Plateau crustal flow pattern. Illustrated in Fig. 6a is a map of $Q_{Lg}(0.2–2.0\ Hz)$, along with the main fault systems (light-blue lines) and geo-tectonics (white lines). Red contours in Fig. 6a delimit the heavily attenuated regions with $Q_{Lg}(0.2–2.0\ Hz) < 200$. A low-$Q_{Lg}$ belt is distributed along the northern and eastern borders of the Tibetan Plateau. Based on this pattern, the main flow channel appears starting from the north and moves eastward. Then, the channel turns toward southeastern Tibet, moving along the western edge of the rigid Sichuan basin. After passing the narrow channel, it spreads to a wide front. In addition to the main flow channel, there may be another channel, which starts from southern Tibet, crosses the Eastern Himalayan syntaxis, and merges with the main channel.

Shown in Fig. 6b is the Pn velocity in the Tibetan Plateau and surrounding areas (Liang and Song, 2006). Comparing Figs. 6a and 6b, the $Lg$ wave attenuation and Pn velocity are generally correlated, with high attenuations being related to low Pn velocities. Normally, the $Q_{Lg}$ reflects the material properties in the crust, while the Pn velocity is related to the properties in the uppermost mantle. Investigating the distributions of both $Q_{Lg}$ and the Pn
velocity anomalies can provide information on the depth dependence of subsurface processes. For the main flow channel, both strong Lg-wave attenuation and low Pn-velocity anomalies are observed, suggesting it happened at greater depth. However, for the second flow channel, although there are strong Lg-wave attenuations, no Pn-velocity anomaly is observed, suggesting it occurred at a much shallower depth. These findings are consistent with the previous observations (Beaumont et al., 2004; Clark and Royden, 2000; Fan and Lay, 2003a; Nelson et al., 1996; Owens and Zandt, 1997; Royden et al., 2008; Xie et al., 2004).

Based on the geophysical evidence, Klemperer (2006) suggested the active flow patterns shown in Fig. 6, where black open arrows indicate the middle- or lower-crust flow directions, while the black open circles mark the regions with no flow. The north-south compression and east-west extension of Tibet drive an eastward flow beneath the Qiangtang and Songpan-Ganzi-Hoh xil terranes. This flow bifurcates north and south of the rigid Sichuan basin. The gravitational potential energy and orographic exhumation drive a southward flow between the subducting Indian lower lithosphere and the brittle upper crusts of the Himalaya and southern Lhasa terrane. Bai et al. (2010) produced magnetotelluric images showing two major zones or channels of high electrical conductivity located at depths 20-40 km and extended horizontally for more than 800 km from the Tibetan Plateau to the Yungui Plateau. Using blue arrows, we superimpose their flow model on both the Q_{lg} and the Pn velocity maps in Fig. 6. The first high conductivity channel is roughly consistent with the regions having both strong attenuation and low Pn-velocity anomalies. Comparing crustal flow patterns proposed by Bai et al. (2010) based on their magnetotelluric data (blue arrows in Fig. 6) with the flow channels constrained by our Q_{lg} data, both of them share a similar flow trend
although their locations are not exactly correlated. In their magnetotelluric observations crossing the eastward low-\(Q_{Lg}\) branch, Bai et al. (2010) did not find any high conductivity layer. However, their magnetotelluric observation did find a flow channel, which stretching from southern to southeastern Tibet, moving clockwise around the eastern Himalayan. This flow channel is consistent with the regional structural trend (e.g., Molnar and Lyoncaen, 1989) and roughly correlates to our second flow channel.

5. Discussion

The \(Q_{Lg}\) variation can result from the physical properties and thermal status of the crust, or be affected by the geometrical parameters of the crustal waveguide such as the change of the crust thickness (Zhang and Lay, 1995). Zhao et al. (2010, 2013) investigated the relationship between the low-frequency \(Q_{Lg}\) and the crust thickness in Northeast China and North China Craton, and found a tendency that the \(Q_{Lg}\) is usually high for regions with thicker crust. However, the Tibetan plateau is characterized by unusually thick crust and very low \(Q_{Lg}\) making it unique and should not be fit into a conventional crust model. The circles in Fig. 7 illustrate the \(Q_{Lg}\) (0.2 – 1.0 Hz) versus the crust thickness for geo-blocks in the Tibetan Plateau and its vicinity. These low-frequency \(Q_{Lg}\) values are averaged for individual geology units and over the frequency band between 0.2 and 1.0 Hz. The average crust thickness data are calculated from the 2\(^\circ\)×2\(^\circ\) global crust model CRUST2.0 (Bassin et al., 2000). To compare the Tibetan data with those from other regions (Zhao et al., 2010, 2013), we include the \(Q_{Lg}\) (0.2 – 1.0 Hz) from the Northeast China (triangles) and North China Craton (squares) into this figure. Both NEC and NCC data are averaged values from large regions rather than from
small geology units, thus are relatively stable and representative. The NEC and NCC data are mostly located at the upper-left in the figure, having relatively thin crust and high $Q_Lg$ values and forming a rough relationship with high $Q_Lg$ values correlate to thicker crust (as shown by a dashed line). On the contrary, the data from the major part of the Tibetan Plateau (circled by the dashed line) are characterized by very thick crust of 57 to 67 km and low $Q_Lg(0.2 – 1.0\ Hz)$ from 163 to 254. The group of data circled by the dotted line comes from stable basins and rigid mountain areas surrounding the Tibetan Plateau (refer to Fig. 1). They have average crust thicknesses between 40 and 52 km and $Q_Lg(0.2 – 1.0\ Hz)$ from 292 to 363. The shaded symbols mark the centroids of different data groups, where triangle and square are averages for the entire NEC and entire NCC, and the two shaded circles are for the part of Tibet with elevations large than 4000 m and the average from stable regions surrounding the plateau.

Fig. 7 suggests that the dominant mechanisms of the Lg wave attenuation in the major Tibetan Plateau differ from those in other regions. Possibly, the strong thermal activities and heterogeneities from partially melted magma chambers make both intrinsic and scattering attenuations very effective. At the same time, bumpy topography, weaker Moho discontinuity and possibly the double-layered thick crust are all unfavorable for the propagation of Lg as a guided wave. The data from stable regions surrounding the plateau have thicker crust and slightly lower $Q_Lg$ compared to the NEC and NCC data. However, their crust thickness and $Q_Lg$ are still different from those in the major Tibetan Plateau. There are two exceptional blocks, the YP and QM, which cannot be assigned to either group. Compared to other regions surrounding the major Tibetan Plateau, the YP and QM have similar crust thickness but their $Q_Lg$ are about 100 lower. What causes this is still not fully understand but the YP and QM are
two weakest areas surrounding the major Tibetan Plateau. Due to the combined effect of the subducting India Plate and strong holding of the Eurasia Plate at TB, QB, OB and SB, the crust and upper mantle material in Tibetan Plateau could be streamed out in between these rigid basins. It has been suggested by previous authors that the YP and QM are two possible paths of material out flow (e.g., Wang et al., 2013; Zhang et al., 2011). It is possible that the crusts in these narrow paths underwent strong shear movement and are highly heterogeneous, which causes additional attenuation.

Fig. 8 summarizes the geology and attenuation in a north-south profile along longitude 93°E. Shown in Figs. 8a and 8b are cartoons depicting the surface topography with geological sutures as well as the crust and upper mantle structures adapted from Jin et al. (1996), Owens and Zandt (1997), Kosarev et al. (1999), Kind et al. (2002), and DeCelles et al. (2002). The broadband $Q_{Lg}$ versus frequency are shown in Fig. 8c (note the vertical coordinate is the frequency). Overlapped on the generally smoothed background attenuation, there are three absorbing bands labeled with $\textcircled{1}$-$\textcircled{3}$. If scattering dominates the $Lg$ wave attenuation, the maximum absorbing happens at $ka = 1$, where $k$ is the wavenumber and $a$ is the dominant scale of the scatters either in the crustal waveguide (Wu et al., 2000) or at the surface (He et al., 2008). Thus, the absorbing band may provide information regarding the scales of heterogeneities. In Fig. 8c, the high-frequency absorbing band $\textcircled{1}$ is located near the northern Himalaya and southern Lhasa terrane. Its frequency band of 2.0 to 10.0 Hz corresponds to small scatters of 60 m to 300 m. Numerous geothermal systems were found in this region (Hochstein and Regenauer-Lieb, 1998). Heat-flow measurements show large variations over a short distance in southernmost Tibet. For example, Francheteau et al. (1984)
reported that the heat flow decreases sharply from 146 to 91 mWm$^{-2}$ over a distance less than 25 km between the southern and northern lakes near the Kangmar domain. The sharp heat flow variations suggest that the regional melting conditions are met at relatively shallow depths in the Tibetan crust. Xie et al. (2004) found a strong crustal Lg-wave attenuation in the Yangbajing graben, and suggested that the attenuation resulted from hydrothermal and magmatic fluids in the upper crust. The absorbing band ② is located in southern Tibet between 28° and 31°N. This area coincides with a series of bright spots found from reflection surveys (Brown et al., 1996; Makovsky et al., 1996; Nelson et al., 1996). These bright spots are located at depths of approximately 15 km and are interpreted to be the top of a low P-velocity layer. Makovsky and Klemperer (1999) used AVO (amplitude versus offset) modeling to constrain the P and S velocities of these spots to be 3.0±0.8 and 1.6±0.8 km/s, respectively, and suggested that the spots were caused by aqueous fluid concentrations that are underlain by partially molten layers. These layers are also characterized by low S velocities (Cotte et al., 1999; Guo et al., 2009; Kind et al., 2002; Rapine et al., 2003), high electromagnetic conductivities (Li et al., 2003; Unsworth et al., 2005), high conductive heat flow (Francheteau et al., 1984; Hochstein and Regenauer-Lieb, 1998) and strong crustal attenuation (Fan and Lay, 2003a; Xie et al., 2004), all suggesting that a minimum-strength layer could exist in the mid-to-lower crust where modern day rheological flow is likely occurring (Klemperer, 2006). The dominate frequency of the strong attenuation is between 0.25 to 2.0 Hz, suggesting heterogeneities of 0.3 km to 2.5 km. The absorbing band ③ spans a distance of approximately 400 km under the Qiangtang and Songpan-Ganzi-Hoh xil terranes. The dominant frequency of this absorbing band is 0.15 to 2.0 Hz, which corresponds to
heterogeneity scales between 0.3 km and 4 km. This is the most volcanically active area in Tibet (e.g., Molnar, 1989; Turner et al., 1993). Owens and Zandt (1997) found a lower crustal low-velocity zone with a high Poisson’s ratio of ~0.35 in this region and suggested that it likely involves partial melt.

The major geology unites in the Tibetan Plateau are aligned in east-west direction and gradually turn to southeast at the east end of the plateau. In Fig. 8d, we roughly project their $Q_{Lg}(0.2-2.0 \text{ Hz})$ values and Pn velocities to the 93°E profile. These average values come from vast areas thus should reflect the behavior of these structures sequentially placed in front of the collision of the Indian Plate. From south to north, the average $Q_{Lg}$ changes by first decrease and then increase. At the southern end, the average $Q_{Lg}$ for the Himalaya Mountains (HM) are around 300. The lowest average $Q_{Lg}$ of 182-185 appear in the Qiangtang (QT) and Songpan-Ganze-Hoh xil terranes (ST). It reaches to 281 in the Qiaodmu basin (QB) and reaches to 397 and 346 in AM and TB. The Pn velocity variation resembles that of the attenuation with its lowest value occurs at QT and ST. However, at northern Tibet, the Pn velocity raises before the increase of the $Q_{Lg}$. If the Pn velocity reflects deeper activities compared to crustal attenuation, Fig. 8d indicates that the shallow activities extended to far north than deeper activities.

6. Conclusion

In this study, we obtained a broadband high-resolution attenuation model in the Tibetan Plateau and its surrounding regions based on regional Lg-wave data. 7545 vertical component seismograms were collected from 146 stations and 232 crustal earthquakes. By visually
checking the 0.5-1.5 Hz bandpassed Lg-wavetrains, we adopted the conventional group
to calculate the Lg-wave spectra. After denoising, both
dual-station and single-station datasets were constructed to jointly invert the Q_{Lg} distribution
and Lg-wave source function. The unevenly distributed sources and stations made the ray
coverage vary geographically. We used the checkerboard method to investigate the inversion
resolution. The best data coverage was in the Tibetan Plateau and between 0.05 and 1.5 Hz,
where the Q_{Lg} model has a high resolution of approximately $1^\circ \times 1^\circ$. Toward higher
frequencies and in the surrounding regions, the resolution deteriorates.

The station amplifications depend on the local rock types, sediment thicknesses and
elevations, etc. We calculated site amplifications and correct them at individual frequencies.

Site responses at 1 Hz are listed in Table S1.

The strong seismic attenuation is one of the prominent characteristics for the crust flow in
the Tibetan plateau (e.g., Klemperer, 2006). The low-Q_{Lg} anomaly can be related to the
changing crustal thickness, sedimentary thickness, high temperature and/or partial melting in
the crust and uppermost mantle (e.g., Fan and Lay, 2003b; Xie et al., 2004; Zhao et al., 2010,
2013). Thus, the low Q_{Lg} values only partially provide evidence for the material flow within
the Tibetan crust. To fully constrain the crustal flow pattern, further measurements such as Pn-,
Sn-, and Rayleigh-wave Q tomography are invoked based on the regional waveform dataset.

Acknowledgments

The broadband data used in this study were retrieved from the China Earthquake
Networks Center (CENC), the Data Management Centre of China National Seismic Network
at the Institute of Geophysics, the China Earthquake Administration (Zheng et al., 2010), the IRIS Data Management Center, and the NEIC. The Pn-velocities used in Fig. 6b were provided by Dr. C. Liang. Some figures were created using the GMT (Wessel and Smith, 1998).

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Figure captions.

Figure 1. A topographic map superimposed with main fault systems (light-blue lines),
regional tectonics in the Tibetan plateau and its surrounding regions, the locations of the
CENC (solid squares) and IRIS (triangles) stations, and the epicenters of selected earthquakes
(crosses) used in this study. The information about the major geo-blocks is also listed in Table
1.

Figure 2. Summary on Lg-wave data processing. As an example, shown here are (a) Raw
Lg-wave spectra (circles) and noise spectra (triangles) recorded at station MC10 from
earthquake 2004/08/26, (b) signal-to-noise ratios (SNR), (c) noise-corrected Lg-wave spectra,
and (d) the numbers of available rays at individual frequencies along with the estimated
resolutions marked by shaded areas. The dashed line in (b) is the threshold used to eliminate
the data with SNR below 2.0.
Figure 3. Selected $Q_{Lg}$ maps compared with previous investigations. (a)-(c) $Q_{Lg}$ maps at 0.5 Hz, 1.0 Hz and 2.0 Hz, (d) a schematic map showing Lg Q measurements from previous investigations, (e) 1 Hz ray coverage, and (f) checkerboard resolution analyses. Also illustrated in the figures are geological boundaries (white or red lines) and the fault systems (thin black lines).

Figure 4. (a) Frequency-dependent $Q_{Lg}$ for the Songpan-Ganzi-Hoh xil terrane (ST), and (b) the $Q_{Lg}$ versus frequency for different geo-blocks.

Figure 5. Selected cross sections of the broadband $Q_{Lg}$. Left column, comparison of average $Q_{Lg}$, surface topography, and Moho depth. Right column, the $Q_{Lg}$ versus frequency. The horizontal coordinate is longitude and latitude is labeled in the figure. Details refer to the text.

Figure 6. Comparison between (a) the average $Q_{Lg}$ (over 0.2 – 2.0 Hz) and (b) the Pn velocity in the Tibetan Plateau and surrounding regions. Also shown in the figure are main fault systems (light-blue lines) and regional tectonics (white lines). Note that the red lines are $Q_{Lg} = 200$ contours which delimit the high attenuation areas. Details refer to the text.

Figure 7. The average $Q_{Lg}$ (over 0.2 – 1.0 Hz) versus the average crustal thickness for selected geology units in the Tibetan Plateau and its vicinity. The data from the major part of the Tibetan Plateau are circled by the dashed line. The data from regions surrounding the
Tibetan Plateau are circled by the dotted line. The triangles and squares are data from the NEC and NCC. Details refer to the text.

Figure 8. A combined longitudinal cross section along 93°E, with (a) surface topography, (b) schematic crust and upper mantle structure, (c) log(\(Q_{Lg}\)) versus frequency, and (d) comparison between average \(Lg\) Q and average \(Pn\) velocity for selected geology units. Details refer to the text.
Highlights

- A high-resolution broadband Lg-wave Q model for Tibetan Plateau.
- Material flow patterns in Tibetan crust constrained by the attenuation model.
- Statistical investigations link the Q model with the plateau dynamics.
- The very low Q and thick crust in Tibet is unique compared with other regions.
- In Tibet, the Lg wave Q and Pn velocity are generally correlated.
Figure 1, Zhao et al., 2013
Figure 2, Zhao et al., 2013
Figure 3, Zhao et al., 2013
Figure 4, Zhao et al., 2013
Figure 5, Zhao et al., 2013
Figure 6, Zhao et al., 2013
Figure 7, Zhao et al., 2013
Figure 8, Zhao et al., 2013
<table>
<thead>
<tr>
<th>Geological block</th>
<th>Block name</th>
<th>CRUST2.0 Crustal thickness (km)</th>
<th>Q₀ (1 Hz Q)</th>
<th>Broadband Q (0.2 - 2.0 Hz)</th>
<th>Low Frequency Q (0.2 - 1.0 Hz)</th>
<th>Pn Velocity (km/s)</th>
<th>Type of blocks</th>
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<tbody>
<tr>
<td>the Tibetan Plateau (&gt; 4 km)</td>
<td>TP</td>
<td>64.1 ± 7.8</td>
<td>280 (194-406)</td>
<td>220 (153-316)</td>
<td>195 ± 41</td>
<td>8.04±0.07</td>
<td>Integrated</td>
</tr>
<tr>
<td>Himalaya</td>
<td>HM</td>
<td>63.6 ± 7.1</td>
<td>386 (315-473)</td>
<td>291 (245-345)</td>
<td>255 ± 68</td>
<td>8.07±0.05</td>
<td>Mountains</td>
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<tr>
<td>Lhasa terrane</td>
<td>LT</td>
<td>65.9 ± 11.4</td>
<td>284 (214-378)</td>
<td>220 (159-303)</td>
<td>196 ± 41</td>
<td>8.03±0.08</td>
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<tr>
<td>Qiangtang terrane</td>
<td>QT</td>
<td>67.0 ± 8.1</td>
<td>238 (171-332)</td>
<td>185 (136-251)</td>
<td>163 ± 35</td>
<td>8.00±0.05</td>
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<tr>
<td>Songpan-Ganzi-Hoh xil terrane</td>
<td>ST</td>
<td>57.8 ± 7.7</td>
<td>217 (147-321)</td>
<td>182 (129-256)</td>
<td>164 ± 24</td>
<td>8.00±0.05</td>
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<tr>
<td>Eastern Kunlunshan terrane</td>
<td>EKT</td>
<td>59.0 ± 4.4</td>
<td>289 (223-374)</td>
<td>218 (160-298)</td>
<td>194 ± 48</td>
<td>8.10±0.06</td>
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<td>Western Kunlunshan terrane</td>
<td>WKT</td>
<td>60.7 ± 5.1</td>
<td>330 (263-414)</td>
<td>259 (212-316)</td>
<td>230 ± 53</td>
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<td>Qaidam basin</td>
<td>QB</td>
<td>57.9 ± 2.5</td>
<td>385 (273-544)</td>
<td>281 (201-395)</td>
<td>250 ± 72</td>
<td>8.16±0.03</td>
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<tr>
<td>Qilianshan mountains</td>
<td>QM</td>
<td>51.4 ± 5.3</td>
<td>315 (252-393)</td>
<td>265 (204-345)</td>
<td>239 ± 37</td>
<td>8.11±0.06</td>
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<td>Alty Tagh mountains</td>
<td>AM</td>
<td>51.8 ± 2.4</td>
<td>517 (439-607)</td>
<td>397 (342-461)</td>
<td>363 ± 83</td>
<td>8.15±0.04</td>
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<tr>
<td>Tarim basin</td>
<td>TB</td>
<td>51.1 ± 3.6</td>
<td>443 (388-506)</td>
<td>346 (301-397)</td>
<td>313 ± 69</td>
<td>8.13±0.06</td>
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<td>Alashan uplift</td>
<td>AU</td>
<td>49.4 ± 3.2</td>
<td>452 (394-518)</td>
<td>343 (304-386)</td>
<td>305 ± 79</td>
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<td>YM</td>
<td>43.0 ± 3.7</td>
<td>444 (385-512)</td>
<td>335 (300-374)</td>
<td>292 ± 76</td>
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<td>OB</td>
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<td>395 (328-476)</td>
<td>328 (283-381)</td>
<td>295 ± 48</td>
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<td>Sichuan basin</td>
<td>SB</td>
<td>40.6 ± 2.3</td>
<td>456 (425-489)</td>
<td>370 (343-398)</td>
<td>343 ± 68</td>
<td>8.06±0.05</td>
<td>Basin</td>
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<tr>
<td>Yungui plateau</td>
<td>YP</td>
<td>43.0 ± 1.2</td>
<td>247 (184-333)</td>
<td>221 (159-305)</td>
<td>203 ± 19</td>
<td>8.02±0.04</td>
<td>Integrated</td>
</tr>
</tbody>
</table>
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